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GRAVITY RECOVERY AND CLIMATE EXPERIMENT

UTCSR Level-2 Processing Standards Document

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DOCUMENT CHANGE RECORD

Issue	Date	Pages	Change Description
01.0	Dec 1, 2003	All	Initial Version
01.1	July 24, 2004	6-16	Corrected TDT and GPS-time definitions in II.1.1; renamed sections II.1.1 and II.1.2; added clarifying text to II.2.1; updated and corrected references
02.0	Oct 5, 2005	All	Updated to RL02 background models
03.0	Feb 19, 2007	All	Updated to RL04 background models
03.1	Feb 27, 2007	8	Corrected typos in values of C30dot and C40 dot within the table in Section II.2.1. Added reference.

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I INTRODUCTION

I. 1 PURPOSE OF THE DOCUMENT

This document serves as a record of the processing standards, models & parameters adopted for the creation of the Level-2 gravity field data products by the GRACE Science Data System component at The University of Texas Center for Space Research (UTCSR). This document is issued once for every release of Level-2 data products generated by UTCSR. That release number is included in the title of this document.

This document applies to the Level-2 GRACE products of the kind (see *Product Specification Document* or *Level-2 User Handbook*)

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This document may be used in conjunction with:

1. GRACE Product Specification Document (327-720)
2. GRACE Level-2 User Handbook (327-734)
3. GRACE GFZ L-2 Processing Standards Document (327-743)
4. GRACE JPL L-2 Processing Standards Document (327-744)
5. GRACE AOD1B Product Description Doc (327-750, GR-GFZ-AOD-0001)

I. 2 DOCUMENT CHANGE HISTORY

This document has been previously issued, in reverse chronological order, for the following Level-2 data product releases:

Product Release	Date Document Issued	Remarks
RL02	November 4, 2005	Doc v.2.0
RL01	July 23, 2004	Doc v 1.1

Several background gravity models were updated since RL02. The estimated parameters have also been updated. ***Please note that UTCSR Release-03 does not exist. This was done in order to synchronize release numbers with other processing centers.***

I. 3 OVERVIEW OF DATA PROCESSING

This section contains a brief overview of the data processing done to obtain the Level-2 products in this release.

The gravity field estimates were made using the conventional dynamic, linear least squares adjustment for the orbit and gravity field from an optimally weighted combination of the GPS & K-Band tracking data collected by the GRACE satellites. Some specifics follow in the next table.

Processing Institution	University of Texas Center for Space Research	
Software Used		
Orbit Software	MSODP	Version 2006.1
Linear Solver Software	AESoP	Version 1.4.5 (beta)
GRACE Data Products Used		
<i>Product ID & Release</i>	<i>Data Rate</i>	<i>Remarks</i>
ACC1B (RL=00 & 01 [†])	1 second	Used in the integrator
SCA1B (RL=00 & 01)	5 second	For observation models & transforming body-fixed accelerations
KBR1B (RL=00 & 01)	5-second Range Rate only	
GPS1B (RL=00 & 01)	30-second Double Differences between one GRACE satellite, one GPS satellite & 2 ground stations	12-station network. IGS Precise Orbits used for GPS satellites – and held fixed during analysis
AOD1B (RL=04)	Used as part of background gravity acceleration models	
Other Notes on Methodology		
Solution obtained as an optimally weighted combination of GPS double differences for each satellite and inter-satellite K-Band Range-Rate – using one-day dynamic arcs over the prescribed data span.		
The statistical constraints or conditioning applied, if any, during the estimation process is described in the README files accompanying the products.		

[†] The RL00 Level-1B data were used between Jan 1, 2003 and Dec 31, 2004; the RL01 data were used for the remainder of the mission.

II ORBIT DYNAMICS MODELS

II. 1 EQUATIONS OF MOTION

The equations of motion for both GRACE satellites are identical in mathematical form. In the remainder of this chapter, the equations will be provided for a single Earth orbiting satellite, with the understanding that the same equations apply to both GRACE satellites. Where appropriate, the parameters or conditions unique to each satellite will be specified.

In the inertial frame

$$\ddot{\vec{r}} = \vec{f}_g + \vec{f}_{ng} + \vec{f}_{emp}$$

where the subscript “g” denotes gravitational accelerations; “ng” denotes the acceleration due to the non-gravitational or skin forces; and “emp” denotes certain empirically modeled forces designed to overcome deficiencies remaining in the force models.

II.1.1 *Time Systems*

The independent variable in the equations of motion is the time system TDT (Terrestrial Dynamical Time). The relationship of this abstract, uniform time scale to other time systems is well defined. The table below shows the relationship between various time systems and the contexts in which they are used.

System	Relations	Notes	Standards
TAI	Fundamental time system	International Atomic Time	n/a
UTC	UTC = TAI – n1 (Time-tag for saving intermediate products)	n1 are the Leap Seconds	Tables from USNO
TDT	TDT = TAI + 32.184s	This is the independent variable for numerical integration.	IAU 1976 Recommendation (equivalent to using TT in <i>IERS(2003)</i> standards)
GPS	GPS = TAI – 19s (time-tag of GRACE observations)	Relationship between GPS & TDT is fixed at 19 seconds	Time-tags in sec since 1200 Jan 01, 2000 GPS Time.

II.1.2 Reference Frames

The fundamental reference frame for the mathematical model is the non-rotating, freely-falling (inertial) reference frame, with the origin defined as the center of mass of the Earth. The Inertial and Earth-fixed reference frames, and their relative orientations and associated standards are further described in the chapter on Earth Kinematics.

II. 2 GRAVITATIONAL ACCELERATIONS

The gravitational accelerations are the sum of direct planetary perturbations and the geopotential perturbations. The vector of direct planetary perturbations is evaluated using the planetary ephemerides. The geopotential itself is represented in a spherical harmonic series with time-variable coefficients, to a specified maximum degree and order, and accelerations are computed by evaluating the Earth-fixed gradient of the geopotential. The accelerations are then rotated (after summation with the non-gravitational accelerations) to inertial frame for the integration of equations of motion. In general,

$$\vec{f}_g = {}_{3 \times 3} M_{ef}^{in}(P, N, R) \vec{f}_g^{ef}$$

The 3x3 rotation matrix M, which depends on Earth orientation is described in the chapter on Earth Kinematics.

Contributions to the spherical harmonic coefficients of the geopotential, and the associated implementation & standards are now compiled. The geopotential at an exterior field point, at time t, is expressed as

$$U_s(r, \varphi, \lambda; t) = \frac{GM_e}{r} + \frac{GM_e}{r} \sum_{l=2}^{N_{\max}} \left(\frac{a_e}{r} \right)^l \sum_{m=0}^l \bar{P}_{lm}(\sin \varphi) [\bar{C}_{lm}(t) \cos m\lambda + \bar{S}_{lm}(t) \sin m\lambda]$$

where r is the geocentric radius, and (φ, λ) are geographic latitude and longitude, respectively, of the field point.

The model used for propagation of the equations of motion of the satellites is called the Background Gravity Model. This concept, and its relation to GRACE estimates, is described further in the *Level-2 User Handbook*. The details of the background gravity model are provided here.

Hereafter, the document *IERS Conventions (2003)* is abbreviated as *IERS-2003*.

II.2.1 Mean Geopotential & Secular Changes

Parameter	Value	Remarks
GM_e	3.986004415E+14	<i>IERS-2003 Standards (value is consistent with using TDT or TT as the time argument)</i>
a_e	6378136.3 m	
$N_{\max}$	complete to degree and order 360	GIF22a is background static model. It uses coefficients of GGM02C (<i>Tapley et al. 2005</i>) from degree 121 to 200, and EGM96 (<i>Lemoine et al. 1998</i>) from degree 201 to 360
$\dot{\bar{C}}_{20}$	+0.116276E-10/y	<i>IERS-2003 (Epoch of C_{20} is 2000.0)</i>
$\dot{\bar{C}}_{30}$	+0.049E-10/y	<i>Epoch 2000.0 (Cheng et al. 1997)</i>
$\dot{\bar{C}}_{40}$	+0.047E-10/y	<i>Epoch 2000.0 (Cheng et al. 1997)</i>
$\dot{\bar{C}}_{21} \text{ \& } \dot{\bar{S}}_{21}$	-0.33745E-11/y +0.16059E-10/y	<i>IERS-2003 (Epoch of C_{21} and S_{21} is 2000.0)</i>
<u>Note 1:</u> The normalization conventions are as defined in <i>IERS-2003</i>, Chapter 6, Eqs 2-3. <u>Note 2:</u> The implementation of computation of spherical harmonics and its derivatives is as described in (<i>Lundberg and Schutz, 1988</i>). <u>Note 3:</u> The degree-1 terms are exactly zero in the geopotential model. <u>Note 4:</u> The mean field GIF22a is an interim product created from a combination of the 22-month time-series of UTCSR Release-02 products. Mean field to degree 120 was adjusted from GRACE data.		

II.2.2 Solid Earth Tides

Solid Earth tidal contributions to the geopotential are computed for the an-elastic Earth model, as specified in Section 6.1, *IERS Conventions (2003)*. Corrections to specific spherical harmonic coefficients are computed and added to the mean field coefficients.

Model	Notes	
Planetary Ephemerides	DE-405	
Frequency Independent Terms (an-elastic Earth)	Degree 2 & 3 – expression in Eq. (1), Ch.6, <i>IERS-2003</i> .	Parameter values from Table 6.1
	Ellipticity contributions from Degree 2 tides to Degree 4 terms	As per Eq. 4, <i>IERS-2003</i>
Frequency Dependent Terms	Corrections to all degree-2 terms	Error in legend for table 6.3a is accounted for (see <i>IERS Conventions Forum</i>)
Permanent Tide in $\bar{C}_{20}$ (zero-tide system)	4.173E-9	Subtracted from total contributions as calculated above (implicitly included in value of the mean C20)

II.2.3 Ocean Tides

The ocean tidal contributions to the geopotential are computed as specified in Chapter 6, *IERS-2003*, Eqs 15. Corrections to specific spherical harmonic coefficients of arbitrary (selectable) degree and order are computed and added to the mean field coefficients.

Model	Description	Notes
Tidal Arguments & Amplitudes/Phases	<i>Doodson (1921)</i> <i>Cartwright & Tayler (1971)</i>	
Diurnal/Semi-Diurnal Band Harmonics	Harmonics of model FES2004 (<i>Lefevre 2005</i>) to degree 100.	Extended to minor constituents by fitting admittances to the provided estimates of the major constituents.
Long-Period Band – all constituents	Self-consistent equilibrium model (<i>Ray 2005 – pers. communication</i>).	

The implementation is as follows. The contributions to C_{lm} and S_{lm} values from all the lines in the Cartwright & Tayler expansion are pre-computed and saved in data files for each calendar day at 10-minute intervals. During orbit processing, the software reads these data files, and interpolates the contributions to the integration or acceleration evaluation epoch. Further details are available in *Bonin 2005*.

II.2.4 Atmosphere & Oceanic Variability

The non-tidal variability in the atmosphere and oceans is removed through using the AOD1B Release-04 product. This product is a combination of the ECMWF operational atmospheric model and a baroclinic ocean model driven with this atmospheric model. The details of this product and its generation are given in the *AOD1B Description Document* (GRACE 327-750, v3, 2007).

This model of the geopotential is available as 6 hourly time series to degree and order 100. The value of the harmonics at intermediate epochs is obtained by interpolation between the bracketing data points.

In order to improve the accuracy of interpolation, the following procedure is adopted. First, the 6-hr epoch values of the S_2 air tide is evaluated from the **Ray/Ponte** model (Ray & Ponte 2003), and subtracted from the 6-hr epoch values of AOD1B product. The remainder is labeled “AOT”, and is interpolated to 10-minute intervals. The **Ray/Ponte** S_2 air tide model is also evaluated at 10-minute intervals and added back to the “AOT” product. Note that, in this way, the values of the resulting product at 6-hr values are identical to the original AOD1B product.

Just as with the ocean tides, the software reads these data files during orbit integration and interpolates the contributions to the acceleration-evaluation epoch.

II.2.5 Solid Earth Pole Tide (Rotational Deformation)

The rotational deformation forces are computed as additions to spherical harmonic coefficients $\bar{C}_{21}$ and $\bar{S}_{21}$, from an elastic Earth model, as specified in Section 6.2, *IERS-2003*.

Model	Description	Notes
An-Elastic Earth Model Contribution to C21 & S21	Scaled difference between epoch pole position and mean pole. See Chapter III (Earth Kinematics) for the linear variation model for the mean pole.	
Polar Motion	Tabular input	<i>IERS C04</i>
Mean Polar Motion & Rates	Linear trend	<i>IERS-2003</i>

II.2.6 Ocean Pole Tide

The self-consistent equilibrium model of **Desai** is used (Desai 2002). A spherical harmonic expansion to degree 100 is used, with the same polar motion time series as for Earth Kinematics or the Solid Earth Pole Tide (See Section II.2.5 or III.1.3).

The contributions to the spherical harmonic coefficients are pre-computed using software provided (http://tai.bipm.org/iers/convupdt/convupdt_c6.html), and stored at 10-minute intervals for each calendar day. The orbit processing software reads these data files and interpolates the contributions to the integration or the acceleration evaluation epoch.

II.2.7 N-Body Perturbations

Unlike the geopotential accelerations, the perturbations due to the Sun, Moon and all the planets are directly computed as accelerations acting on the spacecraft. The direct effects of the objects on the satellite are evaluated using point-mass attraction formulas. The indirect effects due to the acceleration of the Earth by the planets are also modeled as point-mass interactions. However, for the Sun and Moon, the indirect effects include, in addition, the interaction between a point-mass perturbing object and an oblate Earth – the so-called Indirect J2 effect.

Model	Description	Notes
Third-Body Perturbation	Direct & Indirect terms of point-mass 3 rd body perturbations	
Indirect J2 Effect	Sun & Moon only	
Planetary Ephemerides	DE-405	

II.2.8 General Relativistic Perturbations

The general relativistic contributions to the accelerations are computed as specified in Chapter 10 of the *IERS-2003 Conventions*.

II. 3 NON-GRAVITATIONAL ACCELERATIONS

The nominal approach is to use the GRACE accelerometer data to model the non-gravitational accelerations acting on the satellite. The model used is:

$$\vec{f}_{ng} = q \otimes \left[\vec{b} + {}_{3 \times 3} E \vec{f}_{acc} \right]$$

where the q/operator represents rotations to inertial frame using the GRACE Attitude Quaternion SCA1B product; b represents an empirical bias vector; and the 3x3 diagonal matrix E contains the scale factors. The bias vector and scale matrix operate on the GRACE Accelerometer observation ACC1B product, and are estimatable parameters.

II. 4 EMPIRICAL ACCELERATIONS

Empirical accelerations are estimated *from only the GPS tracking data*, as a sinusoidal along-track and cross-track variation with an orbital period (one cycle-per-revolution or

1-cpr). The amplitudes of the sinusoids are estimated as piece-wise constant sine and cosine components over 'sub-arc' intervals of 6-hours.

II. 5 NUMERICAL INTEGRATION

The Predictor-Corrector formulation of the Krogh-Shampine-Gordon, second order, fixed-step, fixed-mesh/order integrator is implemented.

Model	Description	Notes
Dependent Variables		
1. Equations of motion (position/velocity for each satellite)		
2. State Transition Matrix (position/velocity mapping terms only)		
Formulation	Cowell Formulation	
Step-size and Order	5 seconds and 7 th order	

III EARTH & SATELLITE KINEMATICS

III. 1 EARTH ORIENTATION

Earth Orientation here refers to the model for the orientation of the Earth-fixed reference relative to the Inertial reference. The former are necessary for associating observations, models and observatories to the geographic locations; and the latter for dynamics, integration & ephemerides.

Frame	System	Realization
Inertial	ICRS	J2000.0 (<i>IERS-2003</i>)
Earth-fixed	CTRS	IGS2000

The rotation between the Inertial and Earth-fixed frames is implemented as:

$${}_{3 \times 3} M_{trs}^{crs} = Q(t)R(t)W(t)$$

which converts the column array of components of a vector in the terrestrial frame to a column array of its components in the inertial frame. Each component matrix is itself a 3x3 matrix, and is now individually described.

III.1.1 Precession & Nutation

Precession and Nutation are modeled using IAU2000A model (*Capitaine et al. 2002*, *Mathews et al. 2002*). Reference epoch 2000.0 is used. The independent variable is TT since epoch J2000.0 (noon, 01-Jan-2000).

III.1.2 Sidereal Rotation

This rotation is implemented as

$$R = R_3(-GST)$$

where the Greenwich Apparent Sidereal Time (GST) is calculated as the sum of Greenwich Mean Sidereal Time (GMST) and equatorial components of precession and nutation. The GMST calculation uses UT1 as its independent argument, whose evaluation is also summarized in this table.

Quantity	Model	Notes
GMST	Polynomial with UT1 as independent variable.	<i>IERS-2003</i>
Equatorial components of precession & nutation	(<i>Aoki & Kinoshita 1983</i>)	<i>IERS-2003</i>
UT1	cubic interpolation of tabular UT1 corrections	<i>IERS C04</i>
	Diurnal tidal variations adapted from <i>Ray 1995</i> eight constituent model.	<i>IERS(1996) Standards</i>

III.1.3 Polar Motion

The Polar Motion component of rotation is implemented as

$$W = R_1(y_p)R_2(x_p)$$

Quantity	Model	Notes
Tabular variations	Cubic interpolation	<i>IERS C04</i>
Ocean Tidal Variations (Diurnal/Semi-Diurnal)	Orthoweights Formulation	<i>IERS(1996) Standards</i>
Note 1: The rotation matrices are implemented in the small angle, skew-symmetric matrix formulation. Note 2: Rotational deformation accelerations & kinematic station displacements are proportional to the difference between this time-series and a linear model for the pole.		

III. 2 STATION COORDINATES

This section summarizes the models for the mean and time-variable parts of the station coordinates adopted for data processing.

Quantity	Model	Notes
Mean Station Positions and Velocities	ITRF2000	Refers to the position of a geodetic marker or instrument reference point at each site.
Ocean Tidal Loading (Diurnal/Semi-Diurnal band)	Orthoweight Formulation. Orthoweights adjusted to Scherneck model	<i>IERS(1996) Standards</i>
Station Eccentricities	Individual observation models	
Luni-Solar Solid Earth Tidal displacement	Chapter 7, <i>IERS(1996)</i> (Luni-Solar ephemerides from DE-405)	
Rotational Deformation	Scaling of difference of polar motion values from a linear trend model.	See polar motion or Mean Field
Tidal Geocenter (Diurnal/Semi-Diurnal)	Included within Ocean Tidal Loading model	
Atmospheric Loading	Not modeled	
Post-glacial Rebound	Not modeled	
Slow (seasonal) Geocenter Variations	Not modeled	

III. 3 SATELLITE KINEMATICS

The Science Reference Frame (SRF - see *Product Specification Document*) is used in all instances where a satellite-fixed reference frame is needed. The inertial orientation of the spacecraft (i.e. the SRF) is modeled using tabular input quaternion from product SCA1B (*ibid.*). The same product is also used for rotating the accelerometer data to inertial frame prior to numerical integration; for making corrections to the ranging observations due to offset between the satellite center of mass & the antenna location; as well as for computing the non-gravitational forces (if necessary).

At epochs where the GRACE quaternion product is not available, linear interpolation between adjacent values is used.

III.3.1 Rotation of Velocity Components

The position rotations are specified in Section III. 1. The components of the satellite velocity vector are rotated using the matrix approximation

$$\vec{v}_{crs} = M_{crs}^{trs} \vec{v}_{trs} + (PN\dot{R}S) \vec{r}_{trs}$$

III.3.2 GRACE GPS Antenna Offset Model

The GRACE GPS navigation receiver is placed on the top surface (see *Product Specification Document*). For the purposes of orbit and gravity field determination, the antenna phase center location vector for the L3 (LC) ionosphere-free double difference is

$$(0.0, 0.0, -0.419) \text{ meters}$$

in the Science Reference Frame. This is consistent with the value provided in the VGN1B product (RL01). The antenna phase center variation map for the GRACE receiver was not used. The GPS transmitter antenna offsets from (*Haines et al. 2006*) were used.

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